Integrating Geodetic Observations into Hydrological Models

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Land Surface (Hydrological) Models



• Simulate the redistribution of water and energy incident on the land surface using physically based equations

- Merge data from diverse sources, including satellites, in a spatially and temporally continuous and consistent manner
- Economical



System of physical equations:

Surface energy conservation equation Surface water conservation equation Soil water flow: Richards equation Evaporation: Penman-Monteith equation etc.

> Advanced processes such as groundwater storage, carbon fluxes, and vegetation dynamics are beginning to be included

Input and Output Fields

Input Parameters:

vegetation class vegetation greenness/LAI soil type elevation

Required Forcing Fields:

total precipitation convective precipitation downward shortwave radiation downward longwave radiation near surface air temperature near surface specific humidity near surface U wind near surface V wind surface pressure

Summary of Output Fields:

soil moisture in each layer snow water equivalent soil temperature in each layer surface and subsurface runoff evaporation transpiration latent, sensible, and ground heat fluxes snowmelt snowfall and rainfall net shortwave and longwave radiation Mean Root Zone Water Content (%), 1 March 2003



Data Integration Within a Land Surface Model

INTERCOMPARISON and OPTIMAL MERGING of global data fields



Satellite derived meteorological data used as land surface model FORCING

ASSIMILATION of satellite based land surface state fields (snow, soil moisture, surface temp, etc.)



Ground-based observations used to VALIDATE model output





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Integration of Geodetic Observations into Hydrological Models

Challenges

- Observed and modeled variables differ
- Spatial resolution of observations is relatively coarse
- Temporal resolution and latency

Key Advantage

• Geodetic observations are not limited to a certain depth of penetration

Options for Incorporating Geodetic Observations into Hydrological Models



Parameter Definition

Example: Elevation data are used to adjust surface temperature and set the infiltration-to-runoff ratio

Parameter Calibration

Example: The amplitude of GRACE-derived terrestrial water storage variations can help to tune parameters like porosity and depth of soils, for which reliable maps often don't exist

Data Assimilation: Model States

Example: Terrestrial water storage, as derived from GRACE, is merged with the model's own estimate using an optimization algorithm like a Kalman filter

• Data Assimilation: Observed States

Example: Earth's gravity field or rotation is forward modeled based on mass variations simulated by a fully coupled global model, and the GRACE based gravity field is directly assimilated (same theory as radiance assimilation)

• LSMs simulate the terrestrial water cycle, but accuracy is limited by

- quality of the input forcing and parameter data
- model developers' understanding of the physics involved
- simplifications necessary to simulate physical processes economically
- Value of GRACE observations for hydrology is limited by
 - low spatial and temporal resolutions; product latency
 - lack of info on vertical distribution of observed mass changes
- Data assimilation can harness the advantages of each:
 - LSMs provide physically consistent, high resolution output; run up to nearreal time driven by other data
 - GRACE and other observations anchor the results in reality
 - DA incorporates error information to ensure optimal blending







GRACE Data Assimilation Study



- Offline simulations of the Catchment land surface model using GLDAS forcing data
- 10 year spin-up under 2002 forcing
- 20-member ensemble simulations for open loop (OL) and data assimilation (DA)
- Monthly GRACE anomalies: CSR/GFZ/JPL mean, Jan 2003 -May 2006
- Ensemble Kalman smoother DA



three snow layers surface excess root zone excess "catchment deficit"



From scales useful for water

cycle and climate studies...

To scales needed for water resources and agricultural applications

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LDAS models produce continuous time series; near-real time capable.



Models separate snow, soil moisture, and groundwater; GRACE ensures accuracy.



From a global, integrated observation To application-specific water storage components

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GRACE data assimilation improves groundwater storage estimates

	OL		GRACE DA		
	r	<u>RMSE</u>	<u>r</u>	<u>RMSE</u>	<u>skill</u>
Mississippi	0.59	23.5	0.69	18.7	0.20
Ohio-TN	0.78	62.8	0.82	41.1	0.35
Upper Miss.	0.29	42.6	0.29	40.1	0.06
Red-Ark. / L.M.	0.69	30.9	0.72	26.5	0.14
Missouri	0.41	24.5	0.66	19.7	0.20

OL = open loop (no data assimilation)

r = coefficient of correlation

RMSE = root mean square error (mm H_2O)

GRACE Data Assimilation: Recent Progress and Applications

Extension to other regions and the globe



-200 -120 -40 40 120 200 TWS (mm)

GRACE water storage, mm January-December 2003 loop



-200 -120 -40 40 120 200 TWS (mm)

Model assimilated water storage, mm January-December 2003 loop

Arab Land Data Assimilation System project



US and North American drought monitoring project











Summary



• Geodetic observations have certain advantages over conventional ground based and remote sensing observations, thus there is potential for them to contribute to hydrology via model integration

• Potential methods for integrating geodetic observations into hydrological models include:

- Parameter definition and calibration
- Data assimilation: model states or observation states

• Data assimilation has already proven valuable for spatially, temporally, and vertically disaggregating GRACE derived terrestrial water storage variations and mitigating the latency issue

• GRACE data assimilation results are now benefitting water resources applications





Backup slides

Integration of Geodetic Observations into Hydrological Models: Challenges

GRACE

Land Surface Model

Spatial Resolution

Vertical Stratification

Temporal Resolution and Latency

GRACE Data Assimilation

Data assimilation enables information from multiple space and ground based observation systems to be merged in a physically consistent manner, using our knowledge of physical processes as represented in numerical models Catchment LSM (Koster et al., 2000)

Catchment LSM spatial elements (average size ~2,500 km²)





degree of saturation —



GRACE observation scale: river basins (200,000 – 1,000,000 km²)